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Pyrolytic appraisal of the lignin signature in soil humic acids: Assessment of its usefulness as carbon sequestration marker

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ABSTRACT

Lignin markers in humic acids (HA, the alkali-soluble, acid-insoluble soil organic matter fraction) molecular features are explored to assess the extent to which plant biomacromolecules are progressively transformed by humification processes leading to stable C-forms in soils. Humic acids extracted from a collection of mountain calcimorphic soils from Sierra María-Los Vélez Natural Park (Southeastern Spain) under different use and management practices were studied in detail by visible and infrared (FT-IR) spectroscopies and analytical pyrolysis (Py-GC/MS). The HAs display a more or less marked lignin pattern defined by characteristic methoxyphenol assemblages released after pyrolysis that are associated to a typical infrared pattern including absorption frequencies bands at 1510, 1460, 1420, 1270, 1230 and 1030 cm⁻¹. This variability in the HA spectroscopic and pyrolytic patterns was used as a source of molecular-level surrogates to establish the balance between complementary mechanisms of soil C sequestration i.e., a selective preservation of lignin associated to raw organic matter and other plant-inherited macromolecules, or alternative mechanisms involving microbial breakdown or plant precursors and its condensation with microbial metabolites.

We found that HAs in which the lignin signature was comparatively less marked also show high optical density values suggesting unsubstituted, condensed aromatic units and a chaotic organic structure, pointing to the presence of highly resilient carbon forms. Upon analytical pyrolysis, one group of HAs produced major yields of methoxyl-lacking aromatics (alkylbenzenes and alkylphenols), and poor yields of alkyl compounds, which suggest efficient cleavage of biomacromolecules and the occurrence of active microbial synthesis and condensation processes. In fact, these HAs also displayed broadband IR spectra, and visible spectra showing high optical density and polynuclear quinoid chromophores considered of fungal origin. Other group of HAs yielded upon pyrolysis conspicuous series of methoxyphenols and well-defined alkyl series (alkanes, alkenes and fatty acids). The IR spectra also displayed clear lignin and amide bands, as well as intense 2920 cm⁻¹ band and a low optical density, indicative of a marked aliphatic character. This latter is interpreted as the result of recent diagenetic alteration processes of young organic matter and suggests that C sequestration mechanisms in these soils are mainly based on the stabilization of HAs from plant biomacromolecules and aliphatic structures.

These differential lignin alteration patterns indicate that HAs are responsive to soil C sequestration mechanisms, which in the studied soils seem to relay upon microtopographical features rather than to changes in soil use and management.

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1. Introduction

Despite the crucial global implications of carbon (C) stored in soils and sediments [1], the biogeochemical processes involved in C stabilization are not well understood [2]. The study of

the molecular structure and variability in soil organic matter (SOM) may help in unravelling such stabilization processes as well as to infer resilience characteristics of different SOM fractions [3,4]. Although humification is an active process involving biological cleavage of plant and microbial biomass followed by secondary condensation of soluble products into humic substances [5], in some circumstances biodegradation is severely hampered by climatic, biotic or mineralogical soil-forming factors leading to accumulation of raw humus types [6]. In these cases the

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composition of the resulting humic substances could be described as a dynamic heterogeneous mixture of relatively low molecular size components associated via hydrophobic interactions and hydrogen bonds [7]. While at the early humification stages such a supramolecular conformation could be stabilized mainly by weak dispersive forces instead of covalent linkages [8], defining a conceptual model which accounts for their essential role in providing and maintaining soil physical and chemical quality [9], at advanced humification – maturation – stages the humic substances show an outstanding intrinsic resilience and this is crucial to define soil biogeochemical quality of soils temporarily behaving as active C sinks. In particular progressive organo-metallic interactions and additional free-radical condensation of the three-dimensionally bridged structure of humic substances in the course of progressive transformation – maturation – stages may end in comparatively rigid condensed domains. In this situation most of the humic structure could consist of a ‘megamolecule’ formed by a network of C–C and C–O links, where discrete structural units can no longer be recognizable due to the similar stability to chemical and biological degradation of all bonds involved in the whole structure [3]. Such SOM advanced transformation stage is often found in continental Mediterranean semiarid environments where abiotic factors i.e., contrasting temperature and moisture levels, sunlight exposure and the historical effect of wildfires, may lead to soils with low SOM content but highly stable and resistant to biodegradation (resilient).

In this line, several studies have pointed out the possibility to use the structural information provided by the molecular characterization of SOM to differentiate between (a) ecosystems where soil C sequestration relies upon microbial mediated processes with an intense reworking and abiotic condensation of precursors producing intrinsically resilient macromolecular humic substances of chaotic structure and (b) soils where the preservation of raw organic matter prevails and depends on extrinsic factors – mainly organo-mineral interactions – leading to a organic matter organization that is accessible to soil enzymes [10]. Adopting extreme positions about soil C sequestration mechanisms debate have frequently led to hermeneutic controversies in the search of a unified theory justifying SOM stabilization [11]. In particular, calcimorphic soils could be especially suitable to analyze qualitative and quantitative features relevant in the dynamics of SOM. This is due to the fact that these soils displays peculiar features associated both with e.g., microencapsulation processes of particulate plant-inherited materials [12,13] but also with active insolubilization mechanisms of the mineral matrix related to the release of low molecular weight compounds onto a Ca^{2+} saturated soil solution [14]. In fact, the prevailing limestone substrate in semiarid Southern Spain’s soils has been considered to play a role in the low structural variability in the molecular structure of the HAs [12]. This situation demands the use of accurate analytical techniques (i.e., analytical pyrolysis) betraying environmental proxies (molecular markers) responsive of the different sources of environmental variability reflected in the composition of the SOM [15–17].

In the present study analytical pyrolysis (Py-GC/MS) together with visible and IR spectroscopies, are used to study HAs molecular structure and to find compositional descriptors informative about C stabilization processes in a variety of semiarid ecosystems developed on calcimorphic soils.

2. Materials and methods

2.1. Soil sampling

The area of study is located in the Natural Park Sierra María-Los Vélez (Almería, Southeastern Spain) which includes a wide variety of semiarid ecosystems both seminatural (forests and brushwood)

and disturbed (almond-tree orchards and cereal fields) developed on calcimorphic soils. The natural vegetation consists of pine forests (>90 yr), oak forests, reforested pine forests, brushlands, almond-tree orchards and cereal crops (Table 1) [18]. The climate is Mediterranean-type, with typical continental features ranging from semiarid to subhumid. Temperatures are 11.9–16.9 °C with a dry summer season; rainfall events are intense and occasional. The geological substrate consists of sedimentary rocks (limestones, marls and dolomites) and soils are Rendzic and Lithic Leptosols, Calcic, Petrocalcic and Hypercalcic Chernozems, Kastanozems and Hypercalcic, Luvic and Petrocalcic Calcisols [19].

Soil samples (ca. 500 g) were collected with a spade from the uppermost horizon (0–10 cm) after litter removal. In the laboratory, composite samples (obtained by mixing three subsamples taken c. 20 m apart in the field) were air-dried and sieved to fine earth (<2 mm) before analysis.

2.2. Physical and chemical analyses

Soil pH was measured in a 1:1 soil:water suspension. Total carbonates were measured as CaCO_3 with the Bernard calcimeter [20]. The soil water holding capacity was estimated at –1.5 and –0.33 MPa in a pressure-membrane extractor [21]. Total nitrogen was determined by micro-Kjeldahl digestion and soil C by wet oxidation using dichromate in acid medium followed by redox titration [22]. Cation exchange capacity (CEC) and exchangeable ions Ca^{2+} , Mg^{2+} , K^+ , and Na^+ were measured after extraction with ammonium acetate solutions (1 mol L^{−1} NH_4Ac at pH 7) [23].

2.3. Soil organic matter fractionation

The methods applied for the isolation and quantitative determination of the humus fractions were based on standard procedures [14,24]. The separation of the particulate, low density fraction (floating organic particles not yet transformed into humic substances, which in some cases may include some charcoal) was carried out by flotation using soil samples of 10 g suspended in 2 mol L^{−1} H_3PO_4 . After rotary stirring for 1 min, the floating soil fraction or free organic matter was isolated by centrifuging the suspension and filtering, washed with distilled water and analyzed for total C. The soil pellet remaining after centrifugation was resuspended in 0.1 mol L^{−1} $\text{Na}_4\text{P}_2\text{O}_7$ (horizontal motion mechanical shaking for 3 h) and centrifuged. This treatment was repeated up to three times followed by two additional extractions with 0.1 mol L^{−1} NaOH; the dark brown extracts successively obtained (corresponding to the total humic extract: HA + fulvic acid) were aggregated. Two aliquots were taken from this extract, and precipitated with H_2SO_4 (1:1 by vol.) for further determination of the amounts of the acid-soluble fulvic acid and the precipitated HA fraction. The soil residue after the alkaline extraction was washed with distilled water and desiccated at 40 °C. The C concentration in this residue corresponded to the total humin.

2.4. Preparative isolation and purification of the HA fraction

Qualitative isolation and purification (de-ashing) of the HAs was performed by precipitating the total humic extract with 6 mol L^{−1} HCl to pH = 2, centrifuging, redissolving the acid insoluble HA in 0.5 mol L^{−1} NaOH and high-speed centrifuging at 43,500 × g. The centrifugation pellet (particulate organic matter and clay minerals) was discarded and the brown supernatant sodium humate was reprecipitated with HCl and centrifuged. Finally the HA in the gel state and acid pH was dialyzed in distilled water using cellophane bags (Visking® dialysis tubing, molecular weight cutoff 12,000–14,000 Da; pore diameter ca. 25 Å, Medicell) and desiccated at 40 °C.

Table 1
General characteristics of soils in mountain calcimorphic semiarid ecosystems.

Sample/code ^a	Soil type	Vegetation ^b	Geological substrate	Total sand 2–0.02 mm) g kg ⁻¹	Total clay (<0.002 mm) g kg ⁻¹	WHC g cm ⁻³	pH (H ₂ O)	CaCO ₃ (g kg ⁻¹)	Soil C (g kg ⁻¹)	C/N	CEC (cmol _c kg ⁻¹)
2/O	Rendzic Leptosol	Brush encroached	Clastic limestone	391	237	140	8.3	528	18.6	11.5	21.5
3/F	Hypocalcic Calcisol	Relictual oak	Detritic limestones and marls	327	397	283	7.7	278	81.2	18.8	41.9
4/N	Hypercalcic Calcisol	Reforested pine	Detritic limestones	597	189	117	8.0	673	31.3	13.9	25.0
6/I	Hypercalcic Calcisol	Climacic pine	Clastic limestones	414	292	193	8.2	313	32.2	16.1	30.7
7/L	Hypercalcic Petric Calcisol	Cereal	Detritic sediments	572	173	118	8.5	588	13.2	12.4	12.7
8/H	Petric Calcisol	Reafforested pine forest (Pinus halepensis)	Marls and limestones	513	108	59	8.4	940	13.2	16.0	9.2
9/D	Calcic Chernozem	Climacic pine	Detritic limestones	257	272	341	7.9	360	85.4	29.8	53.9
10/B	Calcic Chernozem	Climacic pine	Detritic limestones and dolomies	344	324	267	8.1	544	74.4	37.7	7.9
11/G	Calcaric Rendzic Leptosol	Alpine brush	Limestones	110	490	412	7.7	46	107.4	10.3	50.4
12/M	Gleyic Hypocalcic Calcisol	Orchard	Alluvial calcic marls	69	668	209	8.3	439	6.9	8.1	24.8
14/E	Mollic Calcaric Cambisol	Relictual oak	Limestones	74	554	252	7.6	5	50.3	14.3	37.6
15/J	Hypercalcic Petric Calcisol	Chaparral-like brushland	Limestones	161	475	290	8.0	294	49.0	13.5	35.9
16/P	Calcaric Rendzic Leptosol	Chaparral-like brushland	Limestones	64	516	325	7.9	42	52.9	8.0	37.1
17/K	Hypercalcic Calcisol	Cereal	Limestones	331	293	161	8.2	674	23.4	9.7	17.1
19/A	Mollic Calcaric Cambisol	Climacic pine	Detritic limestone material	508	309	521	7.3	248	195.3	15.1	46.2
20/C	Calcaric Rendic Leptosol	Alpine brush	Limestones	203	482	238	7.9	380	25.0	3.9	26.9

^a One-character code used in Table 3 to refer the contribution of lignin in the HA structure.

^b Reforested pine: *Pinus halepensis*; Bush encroached cleared site: *Stipa tenacissima*, *Lygeum spartum*, *Genista scorpius*, *Artemisia* sp.; Relictual oak: *Quercus ilex* ssp. *rotundifolia*; Orchard: almond trees; Climacic pine: *Pinus halepensis*; Climacic pine: *Pinus nigra* (9, 10), *Pinus halepensis* (19); Alpine brush: *Juniperus oxycedrus*, *Vella spinosa*, *Erinacea antillis*, *Quercus coccifera* (11), *Vella spinosa*, *Erinacea antillis*, *Lygeum spartum* (20); Chaparral-like brushlands: *Quercus ilex* ssp. *rotundifolia*, *Juniperus phoenicia*. WHC: water holding capacity; CEC: cation exchange capacity.

2.5. Characterization of the HA fraction

The HAs were studied by pyrolysis-gas chromatography/mass spectrometry (Py-GC/MS), and by visible and IR spectroscopies. The optical density, that is considered as a surrogate of the aromaticity of HAs [25], was measured in solutions of 66.6 mg L⁻¹ C in 0.02 mol L⁻¹ NaOH. Second-derivative spectra were acquired with a Shimadzu UV-240 OPI-2 spectrophotometer. The IR spectra of the HAs were obtained with a Shimadzu FTIR-8400 PC using KBr dishes with 2.00 mg HA. In order to assist the visual inspection the broad-band profiles in the IR spectra from HAs, a resolution enhancement algorithm was applied based in subtracting the raw spectrum (sized to 640 data points) from multiple (60×) of its second derivative and further application of smoothing [26,27].

Pyrolysis-gas chromatography–mass spectrometry (Py-GC/MS) was carried out using a double-shot pyrolyzer device PY2020iD (Frontier Lab Ltd., Fukushima, Japan) attached to a GC/MS system Agilent 6890 [28]. Samples of 1–2 mg in weight were pyrolyzed in small crucible capsules introduced for 1 min into the micro-furnace preheated at 500 °C. The evolved gasses were directly injected into the GC/MS for analysis. The GC was equipped with a low-to-mid polarity fused silica capillary column DB-1701 (J&W Scientific) (30 m × 250 μm × 0.25 μm film thickness), the oven temperature was held at 50 °C min⁻¹, and then increased up to 100 °C at 30 °C min⁻¹, from 100 to 300 °C at 10 °C min⁻¹, and stabilized at 300 °C for 10 min using a heating rate of 20 °C min⁻¹. The carrier gas used was helium at a flow of 1 mL min⁻¹. The detector was an Agilent 5973 mass selective detector and mass spectra were acquired with a 70 eV ionizing energy. Compound assignment was achieved via single-ion monitoring for various homologous series, via low-resolution mass spectrometry and comparison with published and stored data (NIST and Wiley libraries).

2.6. Statistical analyses

Data treatments were performed with the Statistica package [29]. The joint classification of soil and environmental variables in addition to variables corresponding to molecular characteristics of the SOM was carried out by a multidimensional scaling method [30].

3. Results

3.1. Spectroscopic characterization of the HAs

The HAs chosen for Py-GC/MS studies were selected through a previous exploratory analysis by visible (UV–vis) and infrared (FT-IR) spectroscopy [19]. The results are summarized in Table 2 and Fig. 1 depicting the spectroscopic characteristics of two selected samples showing contrasting levels of residual lignin (i.e., high (19) and weak (16) lignin signature). The high lignin signature sample 19 displays IR conspicuous peaks at 1510, 1460, 1420, 1270, 1230 and 1030 cm⁻¹, just coinciding with a weak intensity of the peaks revealed in the second derivative (as valleys) in the visible spectrum at ca. 620, 570 and 530 nm, typically interpreted as a biomarker feature of specific fungal metabolism [31]. The opposed situation is observed in sample 16, the HA with less marked lignin pattern and high optical density. Due to the suspected importance of the more or less defined lignin signature observed in the resolution-enhanced FT-IR spectra, samples were ranked from high to low similarity to lignin and labelled with an ordinal (IR lignin) used for data treatments as shown in Table 2.

3.2. Pyrolytic characterization of the HAs

Up to 250 different compounds could be detected in the pyrolysates of the studied samples and their relative yields with an indication of the possible precursor compounds [32] is shown in Table 3, where the relative abundance of the groups of compounds is coded with symbols indicating the levels of the semiquantitative percentages related to the total chromatographic area.

Methoxyphenols have been considered as an index for the persistence of plant-derived lignin in the less transformed structural domains of the HAs [33,34]. Thus, accumulation of methoxyphenols would point to comparatively early humification processes with a weak structural alteration of macromolecular lignin, not necessarily requiring its complete breakdown into low molecular weight compounds.

When the HA samples are classified in terms of the more or less marked lignin pattern as seen in the IR spectra, it was possible to classify the samples in a gradient ranging between two extreme categories as previously illustrated in Fig. 1. The group of HAs showing 'weak lignin pattern' was characterized by major peaks of N-containing and carbohydrate-derived pyrolysis compounds, and poor yields of alkyl compounds (e.g., HA samples coded as M,N,O,P (samples #12, 4, 2, and 16), indicative of microbial metabolites accumulation and a more or less efficient biodegradation of plant-derived HA precursors. The major pyrolytic products in these samples consisted of non-methoxylated aromatic compounds (mainly alkylbenzenes), i.e., compounds that have been classically interpreted as typical pyrolytic proxies for matured HAs from terrestrial sources [35]. Despite the low potential of these alkyl-substituted aromatic compounds as source indicator compounds, and to the fact that some might be derived from rearrangements of unsaturated aliphatic compounds during pyrolysis known to occur mainly in the presence of mineral catalysts [36], it is clear that both the stoichiometry of the pyrolysis compounds assemblages as well as the structure of the major fragments in this group of samples (Figs. 2 and 3) clearly differ from those expected from biomass constituents, as it would be the case for soil HAs for which C sequestration mechanisms could be assimilated to a selective preservation of plant and microbial constituents.

On the other side, the HA group showing a 'conspicuous lignin pattern' display a well-preserved methoxyphenol signature (i.e., guaiacol, syringol and their methyl-, ethyl-, vinyl-, propenyl- and acetyl-derivatives) and comparatively lower amounts of alkyl and N-derived pyrolysis products (HAs coded as A,B,C,D,E,F, corresponding to samples # 19, 10, 20, 9, 14 and 3).

Intermediate patterns between these two extreme situations were observed in HAs samples coded as G,H,I,J,K,L, corresponding to HAs # 11, 8, 6, 15, 17, 7, respectively illustrating the above-suggested simultaneous active humification processes contributing to SOM stabilization which also coincides with the occurrence of the mixed vegetation in these sites.

In some cases, the group of HAs with features pointing to a slow biodegradation of the original sources of SOM, i.e., the 'conspicuous lignin pattern', yielded aromatic pyrolysis products of relatively high molecular weight (e.g., biphenyls, compounds 149, 173, 126, 201) and well-defined, wide homologues series of alkanes, α,ω-alkanedienes and fatty acids. An outstanding systematic feature, particularly in HAs from soils developed under natural pine forest, was the release after pyrolysis of polyaromatic compounds (Fig. 2) such as phenanthrenes, retene and diterpene resin acids (e.g., dehydroabietic acid), which coincided with previous pyrolytic descriptors for pine soils in Mediterranean sites [4,37].

The above differences between HA characteristics are in agreement with the previous exploratory analysis using only visible and IR spectroscopies, which also indicated a series of HAs with feature-less IR spectra, high optical density (e.g., at 465 nm) and resolved

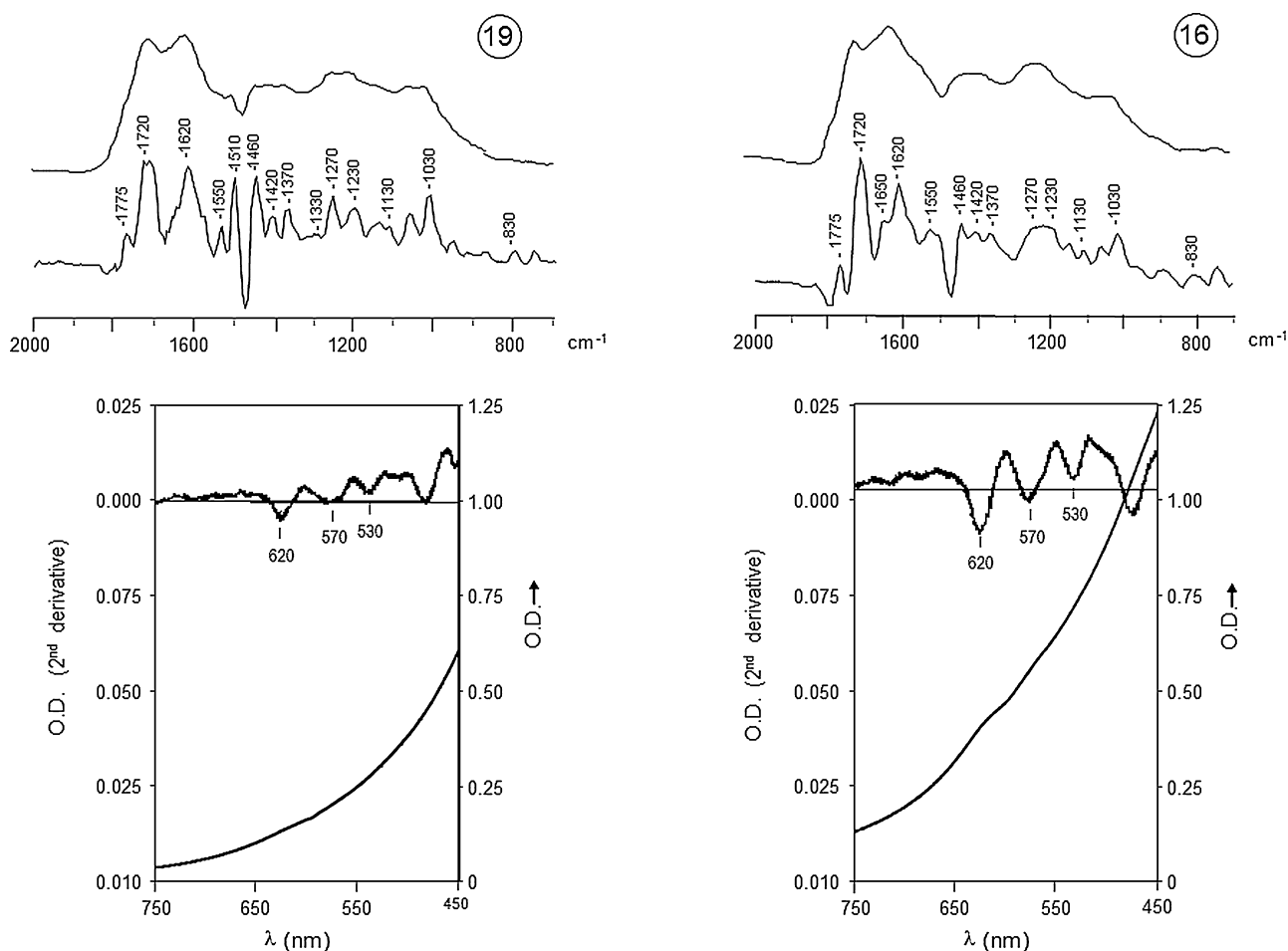


Fig. 1. Spectroscopic patterns of humic acids from semiarid calcimorphic soils representative of extreme levels of residual lignin (left: sample 19 with high lignin signature; right sample 16 with low lignin signature). Top: IR spectra in the 2000–700 cm⁻¹ range, superimposed to the corresponding resolution-enhanced IR spectra. Below: visible spectra and their 2nd derivatives.

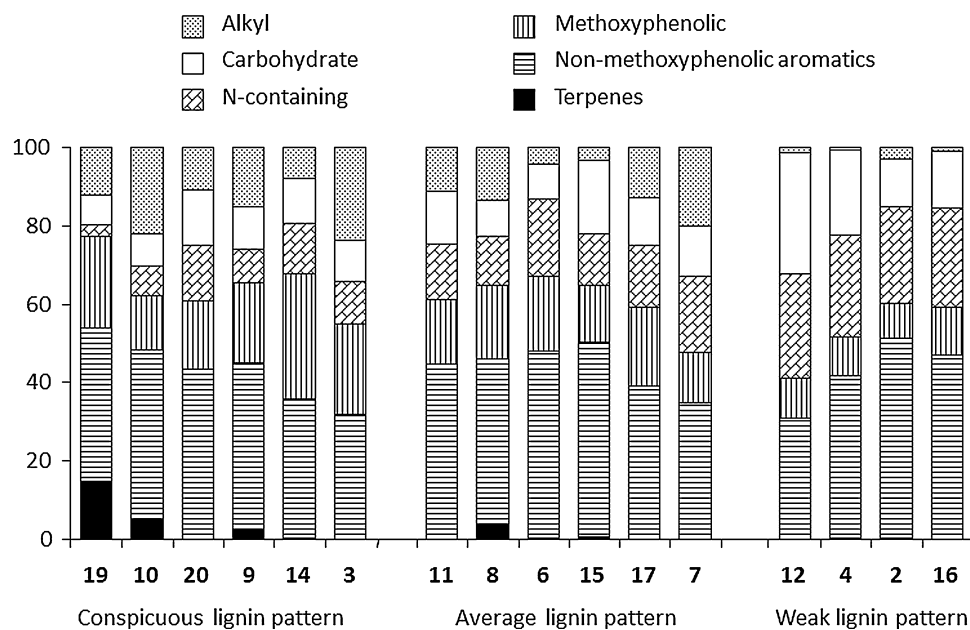


Fig. 2. Relative amounts of different group of pyrolysis products in humic acids from semiarid mountain calcimorphic soils, classified by lignin pattern group as originally defined by the intensity of the lignin pattern in the IR resolution-enhanced IR spectra (Fig. 1).

Table 2

Soil organic matter characteristics and spectroscopic properties of humic acids from mountain calcimorphic semiarid soils.

Sample/code ^a	FOM (C g 100 g soil ⁻¹)	HA (C g 100 g soil ⁻¹)	E ₄₆₅	E 620	IR lignin
2/O	1.2	21.3	1.23	0.0024	5
3/F	2.9	25.5	0.56	0.0016	14
4/N	0.6	54.3	0.66	0.0018	3
6/I	2.0	57.7	0.71	0.0019	13
7/L	2.5	52.0	0.97	0.0018	12
8/H	2.4	51.9	0.63	0.0021	11
9/D	1.6	10.0	0.86	0.0020	10
10/B	1.0	10.5	0.73	0.0019	9
11/G	6.5	89.7	0.84	0.0020	6
12/M	0.9	63.2	1.26	0.0020	2
14/E	4.3	23.8	0.62	0.0009	8
15/J	4.3	12.2	0.95	0.0021	4
16/P	1.4	2.7	1.10	0.0019	1
17/K	1.5	37.6	0.87	0.0021	7
19/A	4.0	13.5	0.52	0.0015	15
20/C	2.5	7.4	0.64	0.0028	16

FOM, free organic matter; HA, humic acid; E₄₆₅, optical density values at 625 nm in the visible range in absorption units; E 620, intensity of the valley at 620 nm in the second derivative spectrum in absorption units; IR lignin, relative intensity of the lignin pattern in the IR spectra.

^a One-character code used in Table 3 to refer the contribution of lignin in the HA structure.

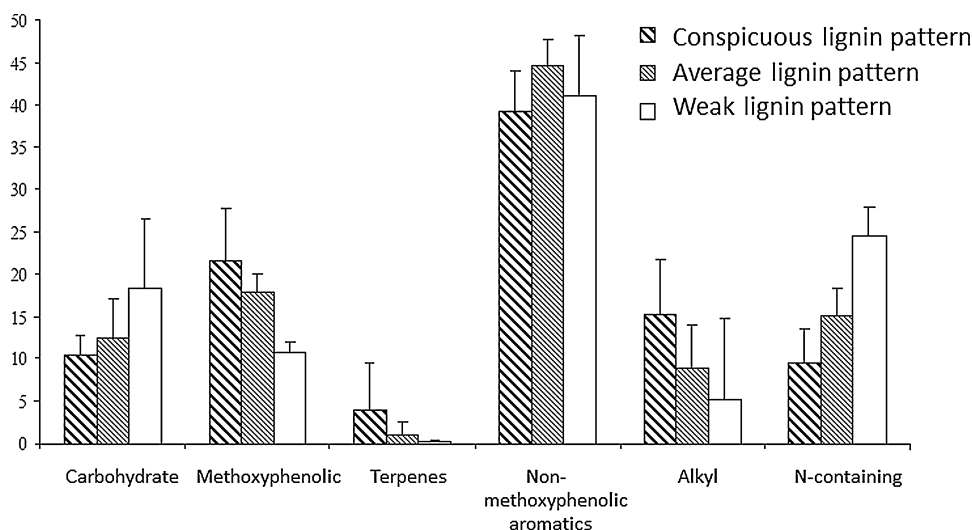


Fig. 3. Average values and confidence limits of relative amounts of the pyrolysis products in humic acids from semiarid mountain calcimorphic soils classified by compound group.

peaks in the 2nd derivative visible spectra characteristic of polynuclear chromophors of fungal origin [31], which coincided with the 'weak lignin pattern' group of samples (Fig. 2). Another group of HAs presenting a comparatively marked aliphatic character and low molecular weight (intense 2920 cm⁻¹ IR band, low optical density and IR spectra displaying typical lignin and amide bands), could be ascribed to group of samples with the 'conspicuous methoxyphenolic pattern' betrayed by analytical pyrolysis.

4. Discussion

The results from Py-GC/MS suggested a variety of structural features in active biogeochemical scenarios which could be ordered as a gradient of progressive humification or SOM maturation [38]. This could be operatively interpreted as the results of a balance between non-excludent soil C sequestration processes, i.e., HAs with lignin domains at early alteration stages vs. HAs formed after heavy structural rearrangement of humic precursors not necessarily with a macromolecular nature.

In fact, in one extreme situation, soil HAs with spectroscopic and Py-GC/MS patterns indicative of heterogeneous composition suggest efficient microbial reworking of the SOM precursors, through humification mechanisms associating products from both plant or

microbial synthesis, as well as secondary molecules and mixtures of oligomers released during litter biodegradation. The Ca²⁺-saturated medium would favour the insolubilization of these humic precursors to be progressively arranged into supramacromolecular mixtures. The broadband IR spectra of these HAs suggest few repeating structural units, and its high optical density is compatible with an advanced diagenetic 'maturation', i.e., selective biodegradation of labile HA moieties and free-radical-induced cross-linking between components that are being randomly incorporated into the HA systems [3].

In other extreme situation, the characteristics of the HAs point to the preservation of comparatively young, lignoprotein-type substances, as betrayed by correlated spectroscopic and Py-GC/MS patterns suggesting that diagenetic stabilization of plant biomacromolecules and aliphatic structures are predominant C-sequestration mechanisms in the corresponding soils. Irrespective to the more or less conspicuous lignin signature in the HAs, the sample groups did not significantly differ in terms of classical stoichiometric indices based on the yields of methoxyphenols (e.g., the classical syringyl-to-guaiacyl ratio) which in many cases lead to accurately distinguish between land use or vegetation cover [39]. This is not unexpected in our case due to the large heterogeneity in terms of vegetation, soil use and HAs degree of maturity.

Table 3Pyrolysis products identified in different groups of humic acids^a with an indication of their relative yields and possible precursor (in brackets).

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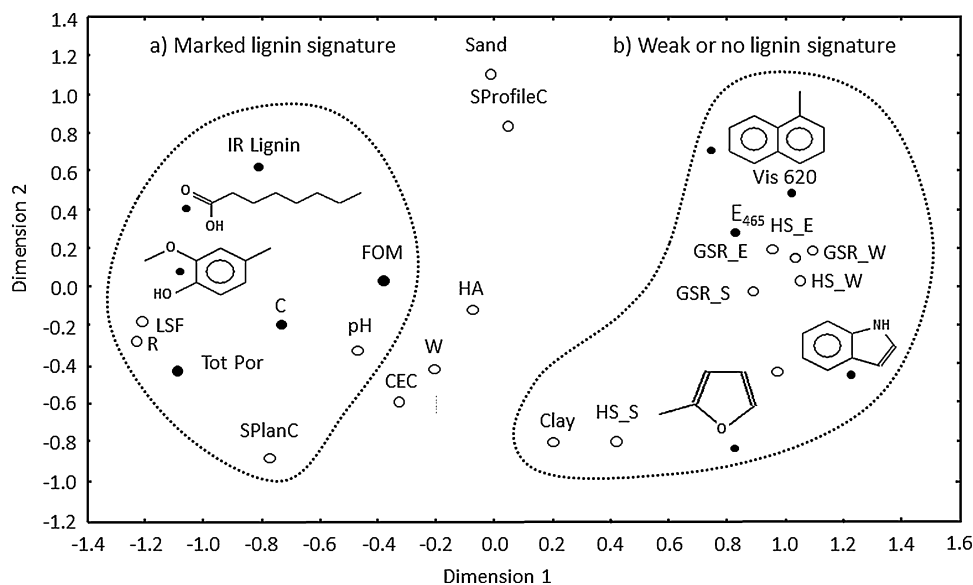


Fig. 4. Automatic classification by multidimensional scaling (final stress = 0.175) of variables corresponding to HA characteristics (solid circles) together with soil analytical properties and environmental factors with a bearing on the humification process (void circles), using the 1-Pearson index as similarity criterion. The plot shows two extreme situations of variable characteristics: (a) humic acids presumably derived from diagenetic transformation of lignins preserved in soil; (b) variables associated to HAs in advanced transformation stages where lignin signature is not evident. Variables in the middle of the scatterdiagram show no strong trends as regards transformation of lignin in the studied soils. Variables obtained with a Geographical Information System: GSR_E, global solar radiation at equinox; GSR_S, global solar radiation at summer solstice; GSR_W, global solar radiation at winter solstice; HS_E, hours of insolation at equinox; HS_S, hours of insolation at summer solstice; HS_W, hours of insolation at winter solstice; LSF, length slope factor; R, contributing area; SPlanC, slope plan curvature; SProfileC, slope profile curvature; W, wetness index. Soil physical and chemical variables and humic acid characteristics (solid circles): C, total soil C; E_{465} , optical density of the HA at 465 nm; FOM, concentration of free organic matter; HA, concentration of soil humic acid; IR lignin, intensity of the lignin pattern in the IR spectra; sand, clay, granulometric fractions; Tot Por, Total porosity; Vis 620, absorption at 620 nm in the 2nd derivative visible spectrum. Total abundances of the main groups of pyrolysis compounds are shown with their formulas.

of a similarity index (in this case the 1-Pearson correlation index) between the processed variables.

In this plot it is evident the association between the 'lignin signature in HAs' with the above indicated pyrolytic yields of methoxyphenols and aliphatic compounds, but also with soil physical and microtopographical features classically associated to moisture levels and accumulation of raw SOM (e.g., FOM). The other cluster of variables included characteristics typical of condensed, black-coloured HAs including fungal quinoid metabolites (E_{465} , Vis 620) in addition to non-methoxylated mono- or polycyclic pyrolysis products. These HA characteristics appear correlated with physical and topographical factors pointing to more conspicuous semiarid features in presumably flat, not intensely eroded areas where clay accumulation and seasonal desiccation would presumably contribute to reach advanced maturation of HAs. In fact, these soil-forming factors are classically described as positive for the humification, i.e., formation of recalcitrant HA-clay complexes and periodic desiccation enhancing SOM insolubilization and condensation processes [14].

Additional statistical treatments mainly correspondence- and discriminant analyses (not shown) failed in evidencing a substantial influence of vegetation types, soil use or direct anthropogenic impact, a situation that has been previously described for similar calcimorphic semiarid mountain ecosystems [12] and attributed to the homogenizing or 'buffering' physical and physiochemical effect of limestone (calcium saturation) on the different organic matter forms (precipitation, encapsulation...). In a similar way, all unsupervised automated classification of the HAs from our soils suggested that the continuous 'gradient' between contrasting 'C-sequestration pathways' observed in our data, was quite independent as regards to local features of soil use and vegetation and the major source of variability being local-scale due to geomorphological factors [40] which are not currently recorded in most studies on SOM dynamics, but that could play a relevant role in

most calcium-saturated soils where humification processes are not controlled by wide contrasts in soil reaction.

5. Conclusions

The analysis of the lignin-derived pyrolytic molecular assemblages in HAs suggests a series of surrogate indicators of SOM quality based on the variable influence of soil C stabilization mechanisms. This was the case with the relative yields of methoxyphenols (which in this study paralleled the intensity of the lignin pattern in the IR spectra, and were negatively related to the E_{465} optical density in the visible spectra) but also with the N-containing, carbohydrate-derived and alkyl compounds. In general the above HA features could be useful for rapid discrimination of the prevailing humification processes influencing C-balance in calcimorphic soils. In this study, SOM quality (i.e., its most advanced stages of transformation) could be assimilated to the extent to which the HAs accumulated in the soil differ in its molecular composition as regards to the biomacromolecules from microbial and plant sources. For instance, the incorporation of specific biomarkers such as conifer resin constituents into the HA structure could represent a valid proxy for raw SOM preservation in forest ecosystems with comparatively low biogeochemical performance. On the opposite, conspicuous concentration of perylenequinone fungal chromophors in HAs (as seen in the 2nd derivative visible spectra) would be pointing to the occurrence of intense microbial reworking of SOM leading to chaotic HA structures (broadband IR spectral profiles) and where the lignin spectroscopic signature is no longer evident by spectroscopic or pyrolytic approaches.

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